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Ocean Currents and Climate

On our watery planet the climate is governed largely by the oceans. But for a long time, discussions of climate change did not take the oceans fully into account, simply because very little was known about them. It seemed that the tremendous masses of water could vary only over a geological timescale. Ideas began to change at the start of the 1960s. Studies of clay extracted from the seabed, supported by new theoretical ideas, suggested that ocean current patterns might shift within mere thousands of years. Other studies began to build a picture of the complex and surprisingly fragile circulation of the world-ocean. In the 1980s, evidence from Greenland ice cores, supported by crude computer models, showed that the North Atlantic circulation could switch radically within a century or two (or maybe less). If global warming triggered such a switch, it would wreak dire harm.

Once scientists asked the question—and it was not an obvious question—the answer was obvious: where are the main ingredients of climate? Not in the Earth’s tenuous atmosphere, but in the oceans. The top few meters alone store as much heat energy as the entire atmosphere, and the oceans average 3.7 kilometers deep. Most of the world’s water is there too, of course, and even most of the gases, dissolved in the water.

It was during the 19th century that the significance of these simple facts became clear. The first thing scientists recognized was how winds passing over the oceans brought moisture and warmth to neighboring lands. Those who sought explanations for climate change included sea changes in their long list of possible causes, and some made this the linchpin. For example, in 1897 a geologist pointed out that a deviation of the Gulf Stream, due perhaps to a gradual raising of land, might set off a glacial epoch.¹

Currents like the Gulf Stream were only minor actors in the story. A far grander feature of the Earth’s surface heat circulation was recognized in the 19th century when scientists tracked down the fact that water hauled up from the depths is nearly freezing, everywhere in the world (there’s a story that the inspiration for these studies came from an old practice of ships in the tropics, chilling bottles of wine by dunking them overboard). This must be water that had sunk in Arctic regions and slowly flowed equatorwards along the bottom. It was a reasonable idea, since water would be expected to sink where the winds made it colder and thus denser.

On the other hand, the warm tropical seas would evaporate moisture, which would eventually come down as rain and snow farther north; this would leave the equatorial waters more salty and

¹ Hull (1897), noting that “the increased snowfall which would thus be caused... would tend to intensify the cold,” p. 107; deflection of currents was likewise seen as central in the scheme of Croll (1875).

therefore denser. So wouldn't ocean waters sink in the tropics instead? The question became part of a long-running debate over what mainly drove ocean circulation: was it differences in density, whether due to cold or salt, or was it the steady push of winds?

Around the turn of the century, the versatile American scientist T.C. Chamberlin took up the question as part of his general program of studying causes of climate variations. He estimated that "the battle between temperature and salinity is a close one... no profound change is necessary to turn the balance." Perhaps in earlier geological eras when the poles had been warmer, salty ocean waters had plunged in the tropics and come up near the poles. This reversal of the present circulation, he speculated, could have helped maintain the uniform global warmth seen in the distant past.¹

Chamberlin and a student of his also drew attention to the crucial role of the ocean in regulating the composition of the atmosphere—as one example, there were "eighteen potential atmospheres of carbon dioxide in the ocean." They noted that a warmer ocean would tend to evaporate more of its carbon dioxide gas (CO₂) and also water vapor into the air, whereas a colder ocean would tend to absorb both gases. These were gases that helped keep the Earth warm through the greenhouse effect. So it appeared that if the planet began to warm up or cool down, the oceans might accelerate the tendency by releasing or taking up the gases. Chamberlin and his student recognized, however, that it was no simple matter to calculate just how the oceans might absorb and emit CO₂. It depended not only on temperature and concentration but on complex chains of chemical reactions.²

Through the first half of the 20th century, scientists ignored the intractable chemical complexities, which hardly seemed worth trying to unravel. Most people assumed as a general principle that over the time-spans relevant to human civilization, natural systems automatically regulated the amount of water vapor and other gases in the atmosphere. In particular, if burning fossil fuels added more CO₂, then as some of the gas dissolved in sea water it would modify the concentration of carbonic acid there, in just such a way that the oceans could absorb all the extra gas.³ The view was fixed in a widely read statement by Alfred J. Lotka. Since the oceans hold many times as much CO₂ as the atmosphere, he explained, it seemed obvious that they must eventually swallow up 95% of any new gas, regardless of the details of the chemistry. The argument was roughly correct in principle (there are about 50 carbon atoms dissolved in the oceans for every one in the atmosphere). But Lotka, in tune with the common assumption that a "balance of nature" kept everything stable, had failed to wonder whether the oceans' absorption could keep up with a really rapid production of CO₂. Scientists of the time assumed without

¹ Chamberlin (1906), quote p. 371; Fleming (1998), p. 89.

² Tolman (1899), quote p. 587.

³ E.g., Gregory (1908), p. 348.

much thought that any change in the atmospheric concentration of any gas could happen only over a geological timescale, hundreds of thousands if not millions of years.¹

The circulation of the oceans was likewise pictured as a placid equilibrium. One pioneer later called this the “underlying theology” of a perpetual steady state of circulation. That was what scientists observed, if only because measurements at sea were few and difficult. Oceanographers traced currents simply by throwing bottles into the ocean. It took them a century to work out the general pattern. “The first law of ocean research,” a leader of the field recalled, “was to never waste your assets by occupying the same station twice! And when this law was violated and the results differed, the differences could be attributed to equipment malfunctioning.” The inevitable consequence, he remarked, was “a climatology steady in time.”²

The classic picture of steady-state circulation was laid out in Harald Sverdrup’s definitive textbook of 1942, drawing on the pathbreaking expeditions of the German oceanographic vessel *Meteor* in the 1920s. Sverdrup described, as one item in a list of many ocean features, how cold, dense water sinks near Iceland and Greenland and flows southward in the deeps. To complete the North Atlantic cycle, warm water from the tropics drifts slowly northward near the surface. Winds presumably added a push to this heat-driven cycle, although the effect of trade winds and the like was entirely uncertain. Sverdrup did not remark that the immense volume of warm water drifting northward might be significant for climate. Like all oceanographers of his time, he gave most of his attention to rapid surface currents like the Gulf Stream.³

Through the 1950s, few scientists found much reason or opportunity to study the slow circulations in the depths. Oceanography was a poorly organized field of research. There were only a few oceanographic institutes, jealously isolated from one another, each dominated by one or several forceful personalities. The funds for research at sea were wholly inadequate to the vast subject. The economics of shipping and fishing supported only studies of practical interest such as surface currents, and little data had been gathered about anything else. The field as a whole scarcely looked like solid science. Theories about ocean circulation had what one expert called “a peculiarly dream-like quality.”⁴

¹ Lotka (1924), pp. 222-24; for the state of thinking in the 1950s, see Hutchinson (1954), pp. 383-84.

² “theology”: Munk (2000a), p. 45; similarly Wunsch (1981), p. 342; “first law”: Munk (2000b), p. 1.

³ Sverdrup et al. (1942), pp. 628-29, 635-37, 647, 685; Sverdrup (1957) likewise sees much more heat transport in the Gulf Stream than across the equator, and describes the North Atlantic deep circulation as driven by winds and heat but not salinity.

⁴ Wenk (1972), pp. 38-41. H. Stommel, privately circulated document, 1954, titled, “Why do our ideas about the ocean circulation have such a peculiarly dream-like quality?” referenced in Warren and Wunsch (1981), p. 601.

Nobody could see a way to do much better. Samples pulled up from thousands of feet down had allowed oceanographers to label the main water masses by temperature and salt content. Thus they could see, among much else, that the water that sank near Iceland had made its way along the bottom as far as the South Pacific. Little more could be said. It seemed impossible to actually measure the motion of these enormous, sluggish slabs of water.¹

Oceanographers concentrated on surface winds and currents, far easier to measure and far more important for seafarers than the slow overall circulation. They had not settled the old debate over how much of the general circulation was driven by the winds, and how much by density changes related to temperature and salinity. Oceanographers who attempted to build theoretical models of the circulation gave their greatest attention to the winds, so meaningful to all who went to sea. Besides, as one of them confessed, “the wind-driven models were easier to formulate.”² Although the calculations were primitive, they gave a starting-point. Bit by bit, important features of the ocean circulation were explained. In particular, in the mid 1950s Henry Stommel threw light on some puzzling old observations by calculating the way cold, salty water could sink in only a few small regions of the North Atlantic and creep along the bottom of the oceans, rising as far distant as the Pacific and returning by various routes.³

Change in the Oceans (1950s-1960s)

The 1950s gave oceanography, like many other fields of geophysics, a breakthrough in organization and funding because of two institutions. First came the U.S. Navy’s Office of Naval Research, which naturally took great interest in every aspect of the subject. The ONR liberally dispensed money for all sorts of research projects, imposing some coordination on the isolated research institutions. Second was the International Geophysical Year (IGY) of 1957-58, which further expanded and strengthened ocean research under the leadership of an international committee. Oceanography was a central player in the IGY, for here as in no other field it was undeniable that progress depended on genuine cooperation among nations, setting aside their political rivalries. Aside from the advantages of having many ships on the seas doing coordinated studies at the same time, even a single survey vessel would be hamstrung without access to foreign ports.⁴

From the outset, the scientists who planned the IGY believed that the role of the oceans in climate change was something they should gather data on, if only for the benefit of future researchers. The point was explained magisterially by the influential meteorologist Carl-Gustav Rossby. Considering how temperature was balanced against salt density, he thought it “not unlikely” that the oceanic circulation “must undergo strong and probably rather irregular, slow fluctuations.” Thus over the course of a few centuries vast amounts of heat might be buried in the

¹ Deacon (1957), p. 81.

² Stommel (1987), p. 58.

³ Mills (1998), p. 634; Warren and Wunsch (1981), pp. xvi-xvii, xxi, 12-13.

⁴ Wenk (1972), pp. 49-50; Miles (1981); Weir (2001).

oceans or emerge, perhaps greatly affecting the planet's climate. In sum, putting climate change and oceanography together would generate important questions and fine opportunities for research. Combining these disparate fields would not be easy, however, and not only because it posed a severe intellectual challenge. Oceanographers and meteorologists worked in separate communities; it would take them decades to establish regular communication and cooperation.¹

The first impressive result of the combined approach was published by a meteorologist, Jerome Namias, in 1963. The previous winter had been phenomenally cold and snowy in North America and Europe. Namias argued plausibly that this was caused, paradoxically, by some unusually warm surface water lingering in a region of the North Pacific. By now oceanographers were taking enough measurements at sea to detect such anomalies, and meteorologists were getting a feel for how a patch of warm sea water might change wind patterns across the entire hemisphere. The patch itself had apparently been maintained by an unusual wind pattern that pushed tropical surface waters northward. It was a persuasive example of what Namias called "complexly coupled mechanisms" leading to a "self-perpetuating system." A change in prevailing winds changed the ocean surface temperature, which in turn influenced the prevailing winds, shifting the planet's weather system at least for a while.²

Namias's work attracted no special notice at the time. It was just one of a number of studies that led to the recognition, in the 1970s, that there were ocean-atmosphere feedback oscillations on a timescale of a few years to a few decades. Most important was the "El Niño-Southern Oscillation" (ENSO) in the mid-Pacific. The breakthrough came in 1969 when Jacob Bjerknes presented a persuasive hypothesis for interactions between what had long been known as separate phenomena: the "El Niño" surface temperature changes in the South Pacific Ocean, and the "Southern Oscillation" of pressure changes in the atmosphere above it.³ Bjerknes's study attracted intense interest once scientists recognized that the El Niño events were connected with powerful if temporary climate anomalies around the world, from torrential rains in Peru to droughts in Kansas. (In the 1980s and 1990s, decade-scale irregular cycles were also detected in the North Pacific, North Atlantic, and Arctic Oceans. These oscillations provided some useful insights for the study of longer-term climate change, and a good workout for computer modelers, but they are not covered in this essay.)

Another set of observations meanwhile cast into doubt the old assumption that the world-ocean maintained an unvarying circulation over many thousands of years. In the mid 1950s, oceanographers managed to drill into the floor of the deep sea, extracting long cores of ooze and clay sediment. Analysis of fossil shells in the cores told much about the condition of the sea water when the sediments had been laid down. Although interpretation of the data was tricky, it

¹ Rossby (1956); translated as Rossby (1959), p. 13.

² Namias's phrases here referred to how snow cover on land would add to the cold, but elsewhere he made a point of wind-sea feedbacks. Namias (1963), quotes pp. 6717, 6185; such a wind-ocean interaction was also reported by Bjerknes (1966).

³ Bjerknes (1969), see also Bjerknes (1966).

seemed to say that the temperature could make a giant jump in as little as a thousand years. Wallace (Wally) Broecker, a young geochemist who had been studying climate changes recorded in ancient lake levels and comparing them with ocean data, began to ask whether “the present configuration is a transient one.” Could it change abruptly with serious consequences for climate? Broecker saw no way to tell whether that could really happen, or ever had. The available data on ocean waters could be interpreted well enough using the traditional model of a torpid, steady-state circulation.¹

A supplementary essay describes how scientists got Temperatures from Fossil Shells, a good example of the ingenious oceanographic techniques and the controversies they could engender.

Broecker was well aware of a bold theory about how ocean changes could turn ice ages on and off rapidly. In 1956 two of his senior colleagues, Maurice Ewing and William Donn, had suggested that raising or lowering sea level could do the trick by letting warm water from the Atlantic spill into the Arctic Ocean or shutting it off. They were drawing on the old tradition of hand-waving ideas about how climate might change in response to the opening or closing of straits, which acted like “valves” controlling the ocean currents that warmed or cooled a region.² This tradition had imagined a gradual geological process, with currents responding passively. But now a few people were ready to speculate, if not in scientific articles than in comments to colleagues, about a more sensitive ocean system.

In 1957 Columbus Iselin, director of the Woods Hole Oceanographic Institution, shared with a journalist some of the talk in the air at Woods Hole. It seemed possible, he said, that during warmer past epochs the North Atlantic water had not been cold enough to sink, so the oceans had stopped overturning. That might happen in future if the greenhouse effect warmed the planet past some critical point. At that point the ocean waters would not so readily absorb CO₂ and carry it into the depths. Thus the level of the gas would climb and greenhouse effect warming would accelerate. It was hard to predict the outcome. If the Arctic Ocean became warm enough to lose its cover of ice, so much moisture might evaporate and come down as snow that it would trigger the formation of continental ice sheets. “Are we making a tropical epoch...,” Iselin wondered, or “starting another ice age?”³

Henry Stommel explored the idea of a radical shift more analytically. He sketched a simple model of the oceans as tanks connected by pipes, with circulation driven by differences of density due to both temperature and salinity. Working through the equations turned up critical points. At these points small change in conditions, even a temporary perturbation, could provoke a “jump” between states. The system, Stommel noted demurely, “is inherently fraught with possibilities for

¹ Broecker (1957), p. III-12; Broecker et al. (1960a); Broecker et al. (1960b).

² Ewing and Donn (1956a); “valves” e.g., Humphreys (1940), pp. 623-24.

³ Robert C. Cowen, “Are men changing the Earth’s weather?” *Christian Science Monitor*, Dec. 4, 1957. Iselin did not mention, but was plainly referring to, the Ewing and Donn model.

speculation about climatic change.”¹ Broecker took up the challenge, speculating that “the Earth has two stable modes of operation of the ocean-atmosphere system, glacial and interglacial.” That would explain a puzzle that came up in the mid 1960s from studies of deep-sea cores. It seemed that slight variations in the planet’s orbit had somehow set the timing for major glacial periods. The orbital variations made only minor changes in the sunlight falling at a given point, and something had to be amplifying the effect. Ocean circulation was a leading suspect.² At a conference held in Boulder, Colorado in 1965, where climate specialists put together for the first time a variety of evidence and ideas about how the climate system could lurch into a new state, speculations about modes of ocean circulation came to the fore.

The circulation of the world-ocean was better charted now, thanks to nuclear physics. Since the 1950s important practical concerns had augmented the purely scientific curiosity of oceanographers. Government officials hoped to bury radioactive waste from nuclear reactors in the abyss. Meanwhile fallout from bomb tests was already mixing into the Pacific Ocean, sparking an international outcry and demands to know exactly where the poisons were going. All these demands to study ocean circulation were fulfilled with a new technique which likewise came from nuclear physics. The radioactive isotope carbon-14 could now be measured in the CO₂ dissolved in a volume of sea water. Pull up a canister filled with water from the depths, and the isotope would tell you how many years had passed since that water had been on the surface, absorbing the gas from the atmosphere. A number of groups took up the challenge. They were financed by agencies that oceanographers of the 1930s never dreamed of, ranging from the International Geophysical Year to the U.S. Atomic Energy Commission.

In the lead was a group at Columbia University’s Lamont Geological Observatory, established by Maurice (“Doc”) Ewing in 1947. Isolated amid woods overlooking the Hudson River outside New York City, Lamont scientists were combining geological interests with oceanography and the new radioactive and geochemical techniques in a burst of creative research. The intense interaction between oceanography and radiochemistry might seem surprising, except that a good fraction of the institution’s funding came from government agencies concerned about the fate of fallout from nuclear weapons tests. Some nine-tenths of Lamont’s funding in its first quarter century derived from military contracts. Not all the scientists were well aware of this strong military connection, which was veiled in secrecy, and they could pursue their research with little attention to anything beyond the purely scientific implications.³ Yet the Lamont group was never far from Cold War concerns as they painstakingly measured carbon isotopes in more than a hundred samples drawn from waters around the world.

When the group began this work in 1955, nobody could say whether the oceans took a hundred years to turn over or several thousand, nor just what paths the circulation followed. The pattern of flow turned out to be different from what Sverdrup had supposed. Tracking down the ages of

¹ Stommel (1961), p. 228.

² Broecker (1966), p. 301.

³ Doel (2001).

various water masses showed that water was moving northward across the surface of the Atlantic all the way up from Antarctica. The return flow of cold water underneath went all the way into the middle of the Pacific. Equally significant was the time scale, which turned out to be half a millennium or so (in particular, the deep water of the North Atlantic had been down there an average of 650 years).¹ Other groups using carbon-14 data agreed that on average the ocean waters took at least several hundred years to turn over. Would that suffice to bury greenhouse gases as fast as humankind produced them? The question prompted Roger Revelle at the Scripps Institution of Oceanography in California to take a close look at the chemistry of CO₂ dissolved in sea water. He showed that the uptake was slow: it would take many hundreds of years for the oceans to dispose of the extra gas we added to the atmosphere.

The full story of the crucial discovery that the oceans cannot rapidly absorb CO₂ is given in a supplementary essay on Revelle's Discovery.

Mapping and Modeling the Circulation (1970s)

As scientists came to see that the ocean circulation might be crucial for climate change, not just over geological time but more immediately, more of them paid attention to the problem. A particularly stimulating idea came from Peter Weyl of Oregon State University in the mid 1960s. He noticed that the moist trade winds that cross the Isthmus of Panama and drop rain into the Pacific Ocean carry fresh water out of the Atlantic, leaving behind saltier water. Weyl built on this to develop a theory of the ice ages, involving the way changes of saltiness might affect the formation of sea ice. He did not publish his model until 1968, but he presented the rudiments at the 1965 Boulder conference. The theory would scarcely have been noticed among the many other speculative and idiosyncratic models for climate change, except for a novel insight.

Weyl pointed out that if the North Atlantic around Iceland should become less salty—as might happen if melting ice sheets diluted the upper ocean layer with fresh water—the entire circulation could lurch to a halt. Without the vast drift of tropical waters northward, he suggested, a new glacial period could begin. Others since Chamberlin had speculated that the circulation might stop if global warming somehow made northern surface waters less dense. Now explicit calculations, however crude, made the idea seem worth studying. Just how precariously balanced was the circulation—which was coming to be called the “thermohaline circulation” (from the Greek for *heat* and *salt*)? And how important was it for climate anyway?²

¹ Broecker et al. (1960b); the circulation pattern was mentioned already by Rossby (1956); translated as Rossby (1959), p. 13.

² Weyl (1968), speculating that the “temporary stagnation” of the bottom water would end because of warming by the interior heat of the Earth; the role of glacial meltwater suppressing North Atlantic Deep Water production was also pioneered by Worthington (1968); a neat explanation of the entire circulation in terms of water evaporating from the North Atlantic more than from the cooler North Pacific was indicated by Warren (1983).

As the 1970s began, the picture of large shifts in the ocean-atmosphere system, only hinted at in cores of deep-sea clay, began to get support from studies of ice cores drilled from the Greenland ice cap. That helped stimulate yet more models for the causes of the ice ages. For example, in 1974 Reginald Newell suggested how oceanic ice sheets could help create “the two preferred modes” for the global movement of heat. When sea ice spread widely (say, around Antarctica) it insulated the sea from the frigid air. The water would no longer get cold enough to sink, and the ocean circulation would decrease. As Newell admitted, all this was guesswork and needed much more study, including numerical modeling.¹

By this point scientists recognized that they would never understand climate change at all until they knew how the oceans worked. “We may find that the ocean plays a more important role than the atmosphere in climatic change,” a panel of experts remarked in 1975. They said that should be “a major motivation for the accelerated development of numerical models for the oceanic general circulation.”² Computer ocean models, however, were primitive compared with atmospheric models. The climate models of the 1960s calculated the general circulation of the atmosphere with, in place of oceans, a “swamp”—a mere motionless wet surface. Yet ocean currents were surely a main component of the climate system. In 1969, the leading modeler Syukuro Manabe used crude measurements by oceanographers to estimate that the currents carried roughly as much heat from the tropics to the Arctic as the general circulation of the atmosphere carried. It seemed that something like half the motor of climate was simply absent from the models.³ (In later decades, better data would show this to be an exaggeration, but it was true that energy transfers through the oceans were a crucial part of the climate system.)

Two obstacles kept modelers from handling the oceans in the same way as the atmosphere. First, while meteorologists measured the atmosphere daily in thousands of places, oceanographers had only a scattering of occasional data for the oceans. And second, while atmospheric models could bypass many difficulties by using a simple equation or a single number to stand in for a complex process like a storm, ocean models could not use that trick. For in the seas, analogous processes lasted months or decades, and had to be computed in full detail. Even the fastest computers of the 1970s lacked the capacity to calculate central features of the movements of energy in the ocean system. They could not even handle something as fundamental and apparently simple as the vertical transport of heat from one layer to the next.

Direct observation showed that heat from the atmosphere was absorbed rather quickly by the upper few dozen meters of sea water (the “mixed layer”), but below that the heat penetrated much more slowly into the cold bulk of the oceans. That bare information was enough for some simple models of climate developed in the early 1970s. Modelers pointed out that if anything added heat to the atmosphere, such as the increase of greenhouse gases, much of it would be absorbed into the upper layer of the oceans. While that was warming up, the world’s perception

¹ Newell (1974).

² GARP (1975), pp. 4, 219.

³ Manabe and Bryan (1969); for empirical evidence they cite Sverdrup (1957).

of climate change would be delayed for a few decades. The first generation of atmospheric general circulation models had entirely ignored that, treating the oceans as simply a wet surface in equilibrium. “We may not be given a warning until the CO₂ loading is such that an appreciable climate change is inevitable,” a panel of experts explained in 1979. “The equilibrium warming will eventually occur; it will merely have been postponed.”¹ Beyond such elementary effects, all was obscure. As one senior oceanographer remarked, scientists had no understanding of the physical processes that brought heat into the depths, but only “a set of recipes.” And even the recipes “may be completely wrong.”²

Another great unknown was the interaction between currents like the Gulf Stream and the giant eddies that the currents spun off as they meandered. Evidence of these broad, sluggishly rotating columns of water had turned up during a survey voyage across the North Atlantic in 1960. This was confirmed by an international campaign carried out by six ships and two aircraft in the early 1970s—another example of how studying ocean phenomena needed international cooperation. The survey discovered eddies bigger than Belgium that plowed through the seas for months. What oceanographers had supposed were static differences in the oceans between their sparse measuring-points, had often actually been changes over time, not over space.

The oceanographers had been vaguely aware that when meteorologists built atmospheric models, they included the energy carried by wind eddies as an important factor (physical oceanography, as one practitioner remarked, was “to some extent a mirror of meteorology”). Yet it was astounding to see what prodigious quantities of heat, salt, and kinetic energy the ocean eddies carried. Indeed nearly all the energy in the ocean system was in these middle-sized movements, not in the ocean currents at all.³

Water is not like air, and computers that could handle meteorological computations were far too slow to work through comparable models for this swirling ocean “weather.” As one ocean modeler complained in 1974, “Extensive research efforts have not yet yielded much more than a greater appreciation of the difficulty of these questions.”⁴ To get a handle on the problem, oceanographers had to understand the oceans from top to bottom. But they had little data on the depths—the occasional expeditions, retrieving bottles of water here and there from kilometers

¹ National Academy of Sciences (1979), p. 2. The thermal inertia of the oceans had been noted earlier, e.g., Sawyer (1972), p. 26, remarked that to come into equilibrium “would take of the order of 100 yr, and in consequence the oceans impose a substantial lag on the response of world temperature...”

² Munk (1966), “recipes” quote p. 728. This example of the struggle with vertical mixing includes a hypothesis about tidal effects; recipes for vertical diffusion may be wrong because mixing may actually happen only “in special places”: Munk (1975) (at a 1972 conference).

³ Survey: J. Swallow and J. Crease on the British ship *Aries*. Campaign: the Mid-Ocean Dynamics Experiment (MODE) 1971-1973, promoted especially by H. Stommel. For history, see Wunsch (1981), 358-59, “mirror” p. 343.

⁴ “extensive research”: Schneider and Dickinson (1974), p. 465.

down, were like a few blind men trying to map a vast prairie. Oceanographers liked to remark that we had better maps of the face of the Moon than of the deep sea. After all, there was little economic incentive, nor much military interest either, in studying the colossal slow movements of water, salts, and heat through the abyss.

The decades-long variations as currents and giant eddies sloshed about in ocean basins were on a scale too great not only for computers, but for human lives. A significant ocean change took longer than making a typical doctoral thesis, sometimes longer than an entire career. “This is bad for morale,” as one oceanographer remarked wryly, and an intellectual obstacle besides. How long would it have taken meteorologists to develop ways to predict weather, if they saw only a handful of storms and cold fronts pass through in a lifetime?¹

Stommel worried that oceanographers did not even know how they could start attacking the problem with their current supply of ideas, techniques, and funds. He felt that researchers were spending their time on “tractable side problems... skirting problems of the ocean circulation” as too tough to handle. The only way forward, he said, would be a concerted group effort.² But it would not be easy to persuade people that this would be worth their time and money. “Thinking about the climate is a relatively new business for oceanographers,” a science journalist reported in 1974, “and despite pressure from their meteorological colleagues many believe that global monitoring and modeling of the oceans... is simply beyond the present capacity of the field.”³

The scientists who made such complaints meant to spur action, and action did get underway. The U.S. government and a few other governments began to give oceanography more money and attention. Already in 1968 the *Glomar Challenger* had put to sea to begin a Deep Sea Drilling Project. Technologies for working on the ocean floor at great depths had been developed extensively for commercial purposes such as oil prospecting, and for Cold War missions including the recovery of sunken submarines and lost nuclear weapons. Now these technologies were put to use for scientific oceanography, including some work related to climate. In particular, increasingly accurate methods had been devised to analyze fossil shells as a “thermometer” for past temperatures. The technique was put to use in the “CLIMAP” project, which in 1976 produced maps of sea temperatures at the peak of the last ice age, roughly 20,000 years ago. As expected, the oceans had looked quite different then—cooler overall, and probably with “a more energetic circulation system.”⁴ But other studies, using different markers, found hints that North

¹ Young (2000), p. 166.

² Stommel (1970), p. 1531.

³ Hammond (1974), p. 1147.

⁴ “Climate: Long Range Investigation, Mapping and Prediction.” CLIMAP project members (A. McIntyre et al.) (1976), quote p. 1136; Cline and Hays (1976); the final product was maps (which I have not seen), CLIMAP (1981); note also CLIMAP (1984). The maps were for 18,000 carbon-14 years ago, now estimated at about 21,000 calendar years.

Atlantic waters had sunk less readily during the last ice age.¹ A number of features were hard to explain, challenging computer modelers to reproduce them.

Another major project that the U.S. government funded in the 1970s was GEOSECS (Geochemical Ocean Sections Study), which studied the present ocean circulation. Teams of researchers sampled sea water at many points, and not only for natural carbon-14. Nuclear bomb tests in the late 1950s had spewed radioactive carbon, tritium, and other debris into the atmosphere. The fallout had landed on the surface of the oceans around the world and was gradually being carried into the depths. Thanks to its radioactivity, even the most minute traces could be detected. The bomb fallout “tracers” gave enough information to map accurately, for the first time, all the main features of ocean circulation. “Now that we have the GEOSECS data,” Broecker boasted, “what more can be said on this subject?”²

The improved grasp of ocean circulation came just in time for a problem that was especially troubling oceanographers: exactly how much CO₂ were the oceans currently absorbing? The GEOSECS data helped them win a long debate with other scientists over the global balance of carbon, taking into account how burning fossil fuels and the destruction of forests emitted gases.³ Yet questions remained about just how the masses of water moved about. Only full-scale computer modeling could give answers, using the GEOSECS data as a reality check. Broecker admitted that “At least a decade will pass before a realistic ocean model can be developed.”⁴

Attempts to represent ocean circulation on computers had begun in the late 1960s. In the lead was Kirk Bryan, a Woods Hole oceanographer who had picked up computer modeling from the enthusiasts at the Massachusetts Institute of Technology while working there for a Ph.D. in meteorology. Joseph Smagorinsky’s atmosphere modeling group in Princeton recruited Bryan to add an oceanographic dimension. Here Bryan and a collaborator, Michael Cox, managed to build a numerical model for a highly simplified ocean basin with five levels. Their computer produced a map of numbers that looked roughly like the Atlantic Ocean’s Gulf Stream and equatorial flow. Bryan later recalled that it was not easy to get such work published. For modeling “was looked at with deep suspicion by many of the oceanographic colleagues as... premature.” Most oceanographers were still struggling to map out what was actually in the oceans, and to understand basic processes (like the giant eddies) which computers were nowhere near able to calculate.⁵

¹ Curry and Lohmann (1982); Boyle and Keigwin (1982).

² Broecker (1981), p. 449.

³ A key model, including diffusion by eddies into the deeps, was Oeschger et al. (1975); see also Broecker et al. (1979), q.v. for references to GEOSECS reports by H.G. Ostlund et al., University of Miami; Broecker et al. (1980).

⁴ Broecker et al. (1980), quote p. 582.

⁵ Bryan and Cox (1968). Bryan, interview by Weart, Dec. 1989, AIP.

Nevertheless Bryan pushed ahead, motivated, as he put it, by “the pressing need for a more quantitative understanding of climate.” One example of the tricks he devised was to revise the equations so they did not include vertical movements of the water surface—the mathematical equivalent of clamping a rigid lid on the oceans. Bryan was pretending that the most obvious feature of oceans, their surging waves, did not exist. That scarcely mattered for the slow circulation, and it speeded up computations tremendously. One major influence remained to account for. Winds helped drive the ocean currents that moved heat from the tropics poleward, and the movement of heat in turn was a main influence on climate. So in 1969 Bryan coupled his ocean basin model to Syukuro Manabe’s model of atmospheric circulation. The pair got a recognizable simulation of a slice of climate (if not exactly our own planet’s climate), and signs of a thermohaline ocean circulation.¹

Others now gravitated toward ocean modeling. A research program that had once seemed “a lonely frontier like a camp of the Lewis and Clark Expedition,” as an ocean modeler recalled in 1975, took on “more of the character of a Colorado gold camp.” One reason was breathtaking advances in computer power. Equally important was the extraordinary improvement in oceanographic data, thanks to GEOSECS and other large-scale ocean surveys. These gave for the first time a three-dimensional picture of the actual oceans in motion—a target the modelers could aim at. They constructed plausible models for individual basins like the North Atlantic and Indian Ocean. It was Bryan’s work that “established the paradigm,” as another expert later remarked. By contrast to the wide variety of approaches in atmospheric models, most ocean modelers created “varietal forms” based on similar physical assumptions and numerical methods.²

Now that highly simplified systems had established that the thing could work, modelers forged ahead with more realistic geography. First to plausibly model the entire global ocean was Bryan’s collaborator Cox, using nine levels of ocean and a grid with boxes measuring two degrees of latitude by two of longitude. The calculations used up so much computer time that he could only follow the ocean circulation a few years in its centuries-long progress, but overall the simulated ocean moved somewhat like the real one. Coupling this to a model global atmosphere gave results that had “some of the basic features of the actual climate,” as Manabe and Bryan quietly boasted. Many unrealistic features remained. Modelers had a long way to go before they could calculate ocean circulation well enough to furnish accurate models of climate.³

¹ Manabe and Bryan (1969); Manabe (1969); Bryan (1969a), quote p. 806; “rigid lid”: Bryan (1969b).

² *Frontier*: Reid et al. (1975), Introduction, p. 3; “paradigm... varietal” McWilliams (1996).

³ Cox (1975) (“the most ambitious ocean simulation so far,” according to W.L. Gates, p. 116); published at the same time was a somewhat cruder whole-ocean model, aimed at integration with the Mintz-Arakawa model, Takano (1975); Manabe et al. (1975), quote p. 3; together with Bryan et al. (1975); further landmarks included Manabe et al. (1979); Washington et al. (1980).

Unpleasant Surprises? (1980-1988)

Another handle on the problem was provided by the Deep Sea Drilling Program (which was followed in 1985 by an international Ocean Drilling Program using the American *JOIDES Resolution*, converted from an oil-drilling ship). In expeditions across the seven seas, year after year workers pulled up thick cylinders of clay and ooze, totaling many kilometers. These were stored in “libraries” which any scientist could exploit for climate studies alongside many other topics.

Studies of fossil shells in the cores gave clues about ocean waters in the past, with a striking conclusion. It now seemed beyond doubt that there had been shifts in the North Atlantic particularly around the end of the last ice age some 11,000 years ago—a time geologists on land had long known as the “Younger Dryas” climate shift. The entire pattern of ocean circulation had evidently changed within a couple of thousand years, or perhaps only a few hundred.¹

That resonated with studies by Willi Dansgaard, Hans Oeschger, and others using cores drilled out of the Greenland ice sheet in the early 1980s. Certain periods such as the Younger Dryas had seen very abrupt cooling around the North Atlantic, episodes so striking that they got a name of their own, the “Dansgaard-Oeschger events.” Meanwhile a study of changes in microscopic deep-sea fossil species showed that the cooling had extended clear to the ocean floor. Such studies using microbiology were not given much credence at the time, however. A bit more convincing was a 1983 report, using the geochemistry of isotopes in fossils, with complex evidence pointing to “a dramatic change in ocean circulation” in the last glacial period. The deep waters of the North Atlantic had apparently grown cold and still. Scientists were being gradually pushed to think about dramatic transitions in the circulation of the North Atlantic, or even the entire world-ocean.²

Oeschger was particularly struck by a rapid rise in the atmospheric concentration of CO₂ at the end of the last ice age, which others had recently discovered in ice and deep-sea cores. The vexing problem of how the gas got in and out of the atmosphere had intrigued him ever since 1958, when he had worked with Revelle’s group at Scripps just as they were discovering that greenhouse warming was plausible. Oeschger also understood that a feedback that released more and more of the gas might accelerate the end of an ice age.

¹ Ruddiman and McIntyre (1981a); Boyle and Keigwin (1982) (using Cd as tracer for nutrients); for further refs., see Broecker et al. (1985); later Boyle and Keigwin, using a core from a spot where deposits had built up exceptionally fast, found that “the deep ocean can undergo dramatic changes in its circulation regime” within 500 years, Boyle and Keigwin (1987), p. 36.

² “A basinwide change of deep water occurred,” Schnitker (1979), quote p. 265; Schnitker (1982) speculated about unstable ocean feedback loops; “dramatic change”: Shackleton et al. (1983), p. 242; Rooth (1982) wrote that “catastrophic transitions in the structure of the thermohaline circulation are not only possible, but have probably occurred on many occasions...,” p. 131.

The main reservoir of CO₂ was the oceans, so that was the first place to think about. In 1982 Broecker visited Oeschger's group in Bern, Switzerland, and explained current ideas about the North Atlantic circulation. Broecker also shared an intriguing new idea: the ocean's uptake of CO₂ during an ice age depended on biochemical changes involving the growth and death of plankton. Oeschger reflected that Broecker's biochemical mechanism would take thousands of years to operate, too slow for the rapid changes found in ice cores. Perhaps, he thought, there had been a transition in the ocean from "a relatively stagnant state" to a state where more rapid mixing brought nutrients to the surface which changed the biochemistry.

The stagnant state might have been caused (as some had earlier speculated) by fresh water flowing in as continental ice sheets melted. That would have diluted the surface salt water until it would not sink, halting the circulation. Many questions remained, Oeschger conceded. But major circulation changes might well have been involved—perhaps triggered by some little perturbation.¹

Oeschger was one of the first to worry that a switch between ocean circulation modes might be set off by the greenhouse gases that humanity was adding to the atmosphere. But as a colleague recalled, "his early warnings were often greeted with disbelief." Oeschger tried to find collaborators to write a paper on circulation modes for submission to a top scientific journal like *Nature* or *Science*, but he met only skepticism and gave up the effort. Two colleagues at Bern did publish a paper in *Nature* suggesting that "ocean circulation changes were the essential cause" of the rapid CO₂ variations seen in ice cores, giving Oeschger credit for the idea, but like Broecker they concentrated on biochemical changes rather than the circulation as such. Oeschger continued to bring the idea up at scientific meetings. Broecker heard him, and his interest was stimulated.²

As we have seen, Broecker had been thinking for decades about possible ocean instabilities. The reports of big, rapid CO₂ variations in Greenland ice cores stimulated him to put this interest into conjunction with his oceanographic interests, since nothing but a major change in the oceans could cause such a swift and global shift in the atmosphere. In fact, scientists later realized that the rapid variations seen in the ice cores had been misinterpreted. They did not reflect changes in atmospheric CO₂, but only changes in the ice's acidity due to dust layers (something had indeed changed swiftly, but not necessarily the CO₂ level). No matter: the error had served a good purpose, pushing Broecker to a novel and momentous calculation. Broecker recalled that one day as he sat in Bern, listening to a lecture by Oeschger describing the abrupt variations in his data,

¹ "large-scale circulation changes," Oeschger et al. (1984), p. 303; he cited Broecker (1982b); for meltwater effect he cited Worthington (1968); and in more detail Ruddiman and McIntyre (1981a); in 1990 Broecker cited Oeschger's paper as the first suggestion "that the Greenland events constitute jumps between two modes of operation of the climate system," Broecker et al. (1990).

² Oeschger to Broecker, 11/23/95 and reply 12/4/95, Broecker office files, Lamont-Doherty Geophysical Observatory, Palisades, NY. "Disbelief:" Stocker (1999); Siegenthaler and Wenk (1984).

“an idea hit my brain.... As quick as that, my studies in oceanography and paleoclimatology merged.”¹

In 1985, Broecker and two colleagues published a paper in *Nature* titled, “Does the Ocean-atmosphere System Have More than One Stable Mode of Operation?” Crediting Oeschger as the first to suggest that the apparent CO₂ changes in Greenland ice cores represented a jump between “two modes of ocean-atmosphere-biosphere-cryosphere operation,” the paper continued that “it is tempting to speculate” that Oeschger’s two modes corresponded to different states of the North Atlantic circulation.

Broecker and his collaborators now identified the key: it was what he later described as a “great conveyor belt” of sea water carrying heat northward. Although the GEOSECS survey of radioactive tracers had laid out the gross properties of the circulation a decade earlier, it was only now, as Broecker and others worked through the numbers in enough detail to make crude computational models, that they fully grasped what was happening. They saw that the vast mass of water that gradually creeps northward near the surface of the Atlantic is as important in carrying heat as the familiar and visible Gulf Stream. “It was an easy calculation,” recalled Broecker, “and I was astounded by the amount of heat that it had.” The energy carried to the neighborhood of Iceland was “staggering,” Broecker explained—nearly a third as much as the Sun sheds upon the entire North Atlantic. If something shut down the conveyor belt, climate would change across much of the Northern Hemisphere.

In one sense this was no discovery, but only an extension of an idea that could be traced back to Chamberlin at the start of the century. Few scientific “discoveries” are wholly new ideas. An idea becomes a discovery when it begins to look *real*. Broecker made that happen by providing solid numbers and plausible mechanisms.² Chamberlin had speculated that the circulation could shut down if the North Atlantic surface water became less salty. Now the effect had been calculated. And Broecker pointed out geological evidence that this might actually have happened, at the start of Younger Dryas times. Just then a vast lake dammed up behind the melting North American ice sheet had suddenly drained, releasing a colossal surge of fresh water into the ocean.

¹ “Idea hit:” Broecker (2000), p. 13; Broecker also recalls seeing Oeschger at a 1984 Florida meeting. For further details of Oeschger and Broecker’s work see Broecker and Kunzig (2008), pp. 101-12. On this and faulty data: Broecker, interview by Weart, Nov. 1997, AIP.

² “Astounded”: Broecker, interview by Weart, Nov. 1997, see also Dec. 1997, AIP. Broecker et al. (1985), “jumps... speculate” p. 25; “conveyor belt” and “staggering” heat flow were publicized in Broecker (1987b), p. 87, and laid out fully in Broecker (1991). Sverdrup (1942) thought most of the heat transport northward was in the atmosphere. One early crude estimate indicating considerable ocean heat transport was Oort and Vonder Haar (1976). Arnold Gordon of Lamont-Doherty published the first good description of the “conveyor” in 1986, but regarding heat transport he noted only that “The continuity or vigor of the warm water route is vulnerable to change” and would influence climate: Gordon (1986). For the history see also Manabe and Stouffer (2007).

Broecker's impressive idea was typical of many ideas in geophysics for the way it drew upon several different areas of data and theory. His own career (as may be seen elsewhere in these essays) had rambled through a variety of fields. Ever since the days when he had trudged around fossil lake basins in Nevada for his doctoral thesis, Broecker had been interested in sudden climate shifts. The idea had remained in his mind while he studied the Atlantic Ocean's circulation as revealed by radioactive tracers, the geo-biochemistry of surface sea water as reflected in deep-sea cores, the timing of sea level changes as measured in coral reefs in New Guinea, and numerous other seemingly unrelated topics. "It's like doing a picture puzzle," he remarked. "You get stuck on one, and then it just sits there. And then along comes an idea, and you say, 'Oh my God, that's a piece that fits right there.'" The trick was to keep many pieces on the table, which meant keeping several different lines of research going at the same time.¹ When one piece fitted into another an unexpected picture could appear, like the possibility of a sudden shutdown of the North Atlantic conveyor belt.

The paper by Broecker and his collaborators made a stir among scientists, less for its new ideas than for putting forth in a plausible and dramatic way hypotheses that until then had been hazy and unappreciated. "Until now," the authors wrote, "our thinking about past and future climate changes has been dominated by the assumption that the response to any gradual forcing will be smooth." Even the most elaborate computer models of climate had shown only gradual transitions—but by their very structure that was all they could be expected to show. In the real world, when you push on something steadily it may remain in place for a while, then move with a jerk.

The numerical ocean models of the 1980s were inadequate to explore such a jerk. Even the fastest computers could still scarcely handle the immense number of calculations that even a quite simple model required. Modelers normally began with a static ocean and ran it through a few simulated decades (or if they could get enough computer time, a century or so) of "spin-up" to watch the currents establish themselves. The models did not get through even a single complete cycle of the globe-spanning circulation that interacted with climate change. As a real-world check, scientists also needed to get a much closer look at the details of the fossil climate record. "Unless we intensify research in these areas," Broecker warned, "the major impacts of CO₂ will occur before we are prepared fully to deal with them."²

In 1987, Broecker followed up with an even more provocative *Nature* paper titled, "Unpleasant surprises in the greenhouse?" Here he emphasized the risk that the current buildup of greenhouse gases might set off a catastrophe. "We play Russian roulette with climate," he exclaimed. He issued the same warning in testimony to Congressional committees, in discussions with journalists and in a magazine article.³

¹ Broecker, interview by Weart, Dec. 21, 1997, AIP.

² Broecker et al. (1985), p. 25.

³ Broecker (1987a), p. 123; Broecker (1987a); Broecker (1987b); U.S. Senate, Subcommittee on Environmental Protection, Hearings, Jan. 26-28 1987, pp. 21-23.

A few scientists and the science writers who listened to them began to warn that the ocean circulation might shut down without much warning, making temperatures plunge drastically all around the North Atlantic. London and Berlin are in the same latitude as Labrador, they pointed out, and would be as barren if not for the prevailing winds that pick up heat from the ocean and carry it westward. Only the more attentive members of the public heard the warning at this time, and they could easily ignore it as just another science fiction speculation.

Among the scientists and others who did pay attention to climate change, prospects of a North Atlantic shutdown became one of the most persistent concerns. A few worried that the North Atlantic region was precisely where most of the data on abrupt climate change came from, and where most of the people who thought about it lived. They wondered whether other trigger mechanisms remained to be discovered elsewhere.

“Does the ocean-atmosphere system have more than one stable mode of operation?” Broecker’s question was already on the mind of computer modelers concerned with future climate change. Even before Broecker published his ideas, Kirk Bryan and others had been working up numerical simulations that included changes in ocean salinity as well as wind patterns. What they found was troubling. A 1985 study suggested that if the level of atmospheric CO₂ jumped fourfold, the ocean’s thermohaline conveyor belt circulation could cease altogether.¹ Another study found that even small perturbations could give rise to radically different modes of ocean circulation. In particular, a spurt of fresh water suddenly released from a melting continental ice sheet—the kind of event that some thought might have triggered the Younger Dryas—could switch the circulation pattern in as little as a century.²

These studies were no more than suggestive, for the models of the mid 1980s were still extremely limited. The planet might be represented in the computer as, to take one example, three equal continents and three equal oceans, extending from pole to pole like the segments of a grapefruit, with the oceans all of uniform depth and the continents without mountains. To keep computation time within reason, Bryan had to hold the cloudiness constant, although he knew clouds would interact with climate change in crucial feedbacks. Along with all that, as Bryan remarked, “uncertainties abound concerning the interaction of the ocean circulation and the carbon cycle.”³

Syukuro Manabe and Ron Stouffer developed a coupled atmosphere-ocean model with more realistic geography. As they were varying the CO₂ to see how that might change climate, they made an inadvertent discovery. If they started two computer runs with the same CO₂ level and other overall physical parameters (the “boundary conditions”), but with different random “initial conditions” for the first day’s weather, they could wind up with two radically different but stable

¹ Bryan and Spelman (1985); the question is the title of Broecker et al. (1985).

² Bryan (1986). N.b. this is Frank Bryan, not Kirk. See Broecker et al. (1990). The cause of the event is still uncertain and under study. For full references see "Rapid climate change" essay.

³ Bryan and Spelman (1985), p. 11,687.

states. In one state, the thermohaline conveyor belt was operating. In the other, it wasn't. The model was still packed with unrealistic simplifications, of course. Yet it seemed at least an “intriguing possibility,” as they put it, that global warming might shut down the North Atlantic circulation within the next century or so, with grave implications for regional climates.¹

Most groups still had too little computer power, and too little understanding, to manage full-scale models of both ocean and atmospheric circulation, let alone link them together. They continued to treat the oceans as a passive “swamp” that exchanged moisture with the air but did little else. That forced the model atmosphere to handle all the transport of heat from the tropics into the polar regions, whereas in the real world ocean currents do a good share of the work. And it entirely missed how heat might sink into the ocean depths. The few teams who attempted ocean circulation models had to use highly schematic geography, and they often left out regions near the poles, which brought mathematical troubles where the longitude lines converged to a point.²

After 1988

In 1989 two groups succeeded in coupling a general circulation model of the oceans to a general circulation model of the atmosphere that had realistic geography covering the entire planet. One of the groups also included a crude model of sea ice. The coupled ocean-atmosphere computer models improved rapidly through the 1990s, and gradually took a central role in thinking about climate change. A variety of studies showed it was quite possible that global warming would indeed shut down the North Atlantic circulation eventually, although the process would probably take centuries.³ Confidence in the validity of models increased as some reproduced the striking El Niño oscillations quite well. Still more encouraging, computer specialists managed to reproduce not only the current state of the atmosphere and oceans but also, using the same models without artificial adjustments, the radically different climate that had prevailed at the height of the last ice age.

Despite these triumphs, much remained to be done before anyone could form a clear picture of how the oceans connected to long-term climate change. Perhaps the most vexing of the many difficulties was figuring in the large amount of CO₂ that the ocean's plankton absorbed from the atmosphere. The plankton population depended on the sea surface temperature, and still more on nutrients brought in by rivers, by wind-borne dust, and by the upwelling of ocean currents—all of which could change as climate changed. The plankton's biochemical behavior meanwhile would affect the chemical balance of sea water, which was also crucial for CO₂ uptake or release.

¹ Manabe and Stouffer (1988), p. 841.

² Schlesinger and Mitchell (1987), p. 796; McGuffie and Henderson-Sellers (1997), pp. 55-56; 1980s work is reviewed in Haidvogel and Bryan (1992); Meehl (1992).

³ Two groups: Washington and Meehl (1989) and Stouffer, Manabe and Bryan (1989); the latter included the sea-ice model by Bryan (1969). Shut down: e.g., Wood et al. (1999); see summary: Rahmstorf (1999).

Alongside these intricately indirect effects, scientists gradually learned to worry about a problem that stemmed directly from the rise of atmospheric CO₂. As ever more of the gas dissolved in the oceans, the acidity of the surface water was measurably increasing. This would make it harder for the water to continue to absorb gas (by the mechanism Revelle had reported back in 1957). The acid might eventually dissolve the calcium-carbonate shells of plankton and other creatures important in marine food chains, with uncertain effects on sea water chemistry. Scientists would have to untangle these complexities before they could truly understand how the oceans' uptake of CO₂ would influence the future climate.

Plenty of surprises were still coming from new data. Especially striking were studies in the 1980s that turned up layers of tiny pebbles in North Atlantic deep-sea cores. The debris could have traveled across thousands of kilometers of ocean in only one way: rafted within far-traveling icebergs. Apparently the North American ice sheet had disintegrated at the edges—perhaps in a gigantic surge?—so that great numbers of icebergs had broken off and sailed the North Atlantic as far as Spain. This fitted with speculations about the breakup of Arctic Ocean ice that had been circulating for decades. In 1988, a German graduate student, Hartmut Heinrich, published evidence that the “iceberg armadas” had swarmed across the North Atlantic regularly at particular phases of the glacial cycle.¹ Further studies showed that these “Heinrich events” connected with the more frequent “Dansgaard-Oeschger” periods of cooling, which seemed to have come roughly every 1500 years during the period of glaciation that ended 12,000 years ago. The exact sequence of cause and effect was obscure, but there was evidence of a link to massive surges of the North American ice sheet and changes in the thermohaline circulation.² In any case, it was now certain that catastrophic climate shifts, connected with shifts in ocean circulation, could affect the entire North Atlantic region, and probably other parts of the globe as well.

Whatever had set off the abrupt shifts, they seemed to have been a feature of glacial epochs, not of warmer times like the present. However, many oceanographers suspected that the present climate was not immune.³ Experts looking into the complexities of the North Atlantic system began to think that it might have a variety of possible modes, not just “glacial” and the present stable “interglacial.” A few scientists on the fringes of the community insisted, on flimsy evidence, that the 1500-year glacial-era cycle identified by Dansgaard and Oeschger continued today, overshadowing greenhouse effect global warming. But a different picture was painted by a new ice core from an Antarctic region where snow accumulated fast enough to show what was happening century by century, so the data could be set alongside the Greenland temperatures. The periodic events were, as many scientists had suspected, a global “seesaw” that redistributed heat from one hemisphere to the other. When the circulation had changed so that the North got cooler,

¹ Heinrich (1988); earlier speculations: Mercer (1969); Ruddiman and McIntyre (1981a).

² Bond et al. (1992); Bond et al. (1993); Broecker suggested that when fresh water was brought into the North Atlantic in a million melting icebergs, it might have halted the North Atlantic thermohaline circulation. Broecker et al. (1992). MacAyeal (1993) proposed an influential theory of the North American ice sheet building up and surging at regular intervals.

³ Bauch et al. (2000); Alley (2000), ch. 15.

the South had grown warmer, and vice versa—a far cry from the current warming all around the globe. (Later work showed it was likely that the Younger Dryas, in particular, had involved a shift of circulation that warmed the Southern Hemisphere while cooling the Northern one.)¹

Meanwhile, increasingly in the 1990s and after, the oceanographers' traditional preoccupation with the North Atlantic gave way to a broader perspective. They began to suspect that the tropical oceans could easily be as important as the North Atlantic in rapid climate change. For one thing, it was becoming plain that even ordinary El Niño events in the tropical Pacific seriously affected weather right around the world. For another, new studies showed that equatorial waters had undergone major changes during past ice ages. Climatologists had believed for generations that ice ages had scarcely affected the equatorial jungles. This was now replaced by a view of the globe as a system where every region reacted to changes everywhere else. Suppose, to take just one possibility, a variation of tropical climate altered how winds carried moisture from the Atlantic to the Pacific, altering the salinity? Broecker and co-workers argued that such variations could drive a feedback cycle that might bring “massive and abrupt reorganizations of the ocean-atmosphere system.”² And there were probably other mechanisms, so far barely visible, adding their own complexities.

Only computer models could say which of these ideas might really work, if any. Coupled ocean-atmosphere models posed a severe challenge even for the new supercomputers. Modelers could not directly compute such crucial features as the giant ocean eddies, but had to represent them with a simple set of average parameters. Nevertheless the modelers managed to simulate abrupt shifts of the North Atlantic circulation, confirming that during a glacial period it could shut off and on by itself. A change of circulation also looked only too likely within the next few centuries as global warming took hold. Paradoxically, by cutting off the flow of heat from the tropics that seemed likely to bring severe cooling to regions from Chicago to Moscow.³ Meanwhile several teams stirred up anxieties by announcing they had observed significant changes in the North Atlantic salinity and circulation. Was the circulation already slowing down? Others pointed out

¹Skeptics: Singer and Avery (2007). The redistribution was suggested by Crowley (1992); Stocker (1998) gave the “seesaw” term currency. Antarctic results: EPICA community members (2006), see Jansen et al. (2007), p. 435. The “bipolar see-saw hypothesis” was further strengthened by well-dated evidence of past temperature changes in a core drilled in the South Atlantic, Barker et al. (2009). Younger Dryas: Barrows et al. (2007).

² Broecker and Denton (1989), quote p. 2489; Broecker et al. (1990); for the evaporation cycle Warren (1983); a detailed review is Broecker and Denton (1990); for a more recent review, Broecker (2000).

³ A key eddy parameterization was by Gent and McWilliams (1990); for this history see Manabe and Stouffer (2007), p. 398-99. Manabe and Stouffer (1993) pioneered the demonstration of a transition under future warming. A later model showed a shutdown was especially likely with rapid increase of greenhouse gas emissions, Stocker and Schnitter (1997); see also Broecker (1997); Ganopolski and Rahmstorf (2001) for instability during a glacial period; IPCC (2001), pp. 439-40.

that this might all be part of a normal decades-long cycle (see below). Eventually, surveys found that the system varied so widely from year to year that it might take decades of observations to confirm any definite trend.¹

Magazines, newspapers, television science shows and even a major action movie had popularized the notion that a sudden “shutdown of the Gulf Stream” could plunge us into an ice age. Experts in related fields belatedly began to pay attention, and pointed out that nothing of the kind was likely. For one thing, the planet-spanning thermohaline circulation was not the same thing as the Gulf Stream. That famous surface current was an inevitable result of wind patterns on a rotating planet, and nothing would shut it down. Another thing that seemed obvious in retrospect, but was not widely recognized until 2002, was that even if the North Atlantic circulation did halt, England’s climate would never be as cold as Labrador’s. For Labrador is frigid because it lies downwind from tundra that freezes in winter, whereas England is warmed by prevailing Westerly winds that pick up the ocean’s heat, much of which is simply heat retained from the summer.²

By this time improved computer models were giving fairly consistent results. The ocean circulation was indeed likely to slow down during the 21st century, but more gradually than some had feared. To the extent that a slowdown might come, any cooling it caused around the North Atlantic would be outrun by increasing warmth caused by the greenhouse gases accumulating in the atmosphere. The international panel that studied these issues concluded in 2007 that a dramatic shift of the ocean circulation was “very unlikely” within the 21st century. Still, there might well be a gradual change that would affect, for example, fisheries.³

As for times beyond the 21st century, the experts could say nothing with confidence. Ocean models were not yet good enough to completely rule out the possibility that eventually greenhouse warming would push the circulation system across some threshold and bring a serious shutdown. The most widely used and elaborate models could hardly be expected to show abrupt changes, for they had been *built* to be stable. Indeed these models had difficulty reproducing the sort of climate system jumps that were abundantly clear in the actual geological record of glacial times. And among all the widely varying models, a few did sometimes show abrupt ocean circulation shifts that brought severe climate change in the space of a few decades.⁴

¹ Lazier (1995); Bryden et al. (2005) (the editor of *Nature* removed a question mark at the end of the paper’s title, according to Petr Chylek, Letter to *Physics Today* March 2007, p. 14). Kerr (2006).

² Seager et al. (2002).

³ A typical model showing a roughly 2°C drop in Europe was Vellinga and Wood (2002). A later comparison of 11 coupled atmosphere-ocean models found the circulation “decreases by only 10 to 50% during a 140-year period (as atmospheric CO₂ quadruples), and in no model is there a land cooling anywhere....” Gregory et al. (2005). “Very unlikely:” IPCC (2007c), p. 16.

⁴ By 2006 some models still found that changes in the ocean circulation “are able to produce abrupt climate changes on decadal to centennial time scales,” Randall et al. (2007), p. 641.

Modelers had not yet fully grasped even the current global ocean circulation. Among other shortcomings, their grid boxes were still too large to realistically represent giant eddies or narrow currents like the Gulf Stream. (In the early 2000s a few supercomputers began to probe such details, but most models still had to make do with averages.) It turned out that the ocean system, like most features of climate, was more complex than the first models had supposed. The planet-spanning thermohaline circulation—or “Meridional Overturning Circulation” (MOC) as specialists began to call it in the 1990s—was driven less by saltiness in the North Atlantic than by winds in the Southern Hemisphere, and by even more subtle effects. New data hinted that much of the heat energy moving vertically from layer to layer in the oceans was not transported by some kind of average convection, as the models had assumed, but was moved by tides. Tidal mixing of coastal waters might be as important as saltiness and winds in driving the Meridional Overturning Circulation, which depended as much on the “pull” of water returning to the surface as on the “push” of water sinking in the North Atlantic.¹

Around 2006 a group of scientists published a still stranger idea. It seemed that one way the ocean layers were mixed was through the churning of countless tiny swimming animals in regions where they proliferated. If climate changes affected these marine populations, the biological feedback could have profound effects on the circulation. The upper ocean layers also seemed to mix together more in regions with many storms, and storminess was another thing that might change as the climate changed. Up to this time, modelers had accounted for mixing by a crude expedient, throwing in a constant parameter that matched the average global data for heat transfer. The emerging understanding of how the oceans were stirred by tides and currents flowing over a rough sea floor, by storm winds, and even by plankton, gave computer modelers a beginning—although only a beginning—in calculating just how and where heat moved up and down through the layers.²

There was also new evidence that the North Atlantic Ocean, all on its own, goes through quasi-regular oscillations. This tended to confirm the cycles around 60-80 years long that Dansgaard and others had found for the region. Since the 1920s, meteorologists had been talking about a decades-long cycle of weather patterns around the North Atlantic. Perhaps it was a slow sloshing of water masses that had made temperatures there, more than in many other regions, rise until the 1940s and then dip until the 1970s. Whatever the cause, would a new long-term fluctuation bring another temporary halt in the warming of this crucial region—once again confusing the public

¹ Munk and Wunsch (1998); Egbert and Ray (2000). On the MOC see Toggweiler and Russell (2008). Note also a suggestion that the “North Atlantic Oscillation” was driven by changes in upper atmosphere wind patterns around the entire hemisphere, Wallace and Thompson (2002). See discussion in Keeling (1998), pp. 70-73.

² “Biomixing:” Schiermeier (2007). Mapping global patterns and incorporating the results in models was described as “...a daunting task... requires a large effort, but ... feasible” for yet another important type of mixing, the breaking of the internal waves on the surfaces between layers of different densities, Gregg et al. (2003). Merryfield (2005) reviews recent mixing studies.

about the greenhouse future? Until scientists understood such major effects, and constructed better models, and stopped interrupting one another with surprising new evidence and ideas, the ocean circulation would remain one of the biggest uncertainties in the equation of climate change.

Progress would depend on data, and oceanographers still had sampled only a minute fraction of the world-ocean. Beginning in the 1970s, collaborative projects had mobilized thousands of people from scores of nations. The march of acronyms started under the international Global Atmospheric Research Program (GARP) with regional studies like the groundbreaking GARP Atlantic Tropical Experiment (GATE), carried out in 1974. Next came a Tropical Ocean-Global Atmosphere study (TOGA) that surveyed the equatorial Pacific, inspired by the devastating El Niño of 1982-83. The advent of realistic supercomputer models in the 1980s fostered a more global view of ocean dynamics, by calculating how waters from the North Atlantic circulated all the way to the mid-Pacific and back. To feed the models, there were now satellites (starting with the short-lived SEASAT of 1978) that could measure winds, waves, temperatures, and currents in the remotest reaches of the seas. But the satellites could not measure everything, and what their instruments did measure required “ground truth” observations for checking and calibration. The global approach was embodied in a World Ocean Circulation Experiment (WOCE), planned in the 1980s and carried out in the 1990s by some thirty nations. It was supplemented by a Joint Global Ocean Flux Study (JGOFS) that looked at CO₂ uptake and other ocean chemistry. A still more ambitious international “Argo” program launched 3,000 sensors around the world between 2004 and 2007. Each unit descended as deep as 2,000 meters (6,600 feet), measured temperature and salinity as it drifted on ocean currents, and returned to the surface every ten days to radio the data to passing satellites.¹

Mining old data could also tell many things. One project burrowed through historical records to transcribe literally millions of thermometer readings, assembling a database for the most basic of all climate numbers—the temperatures within the seas. Since the world-ocean absorbs dozens of times more heat than any other component of the climate system, it was here if anywhere that the reality of global warming should be visible. The team found that the heat content of the upper oceans had risen markedly in the second half of the 20th century, in a pattern that neatly matched the “signature” that computer modelers predicted from the greenhouse effect.²

The coupled ocean-atmosphere models were now good enough to pass other key tests. They could reproduce the main features of climate around the planet, and how it had changed over the past century. The models gave reasonably correct pictures of how sea temperatures changed when perturbed by a great volcanic eruption or even an ice age. By the start of the 21st century, the modelers were confident that they could calculate what would happen in future decades as the level of greenhouse gases in the atmosphere climbed: the world would keep getting hotter. But

¹ For a summary to 2001 see Thompson et al. (2001).

² Levitus et al. (2001) (the data compilation was by the NOAA group Levitus headed); Barnett et al. (2001), updated and improved by Levitus et al. (2005).

they were unable to say just how severe the climate changes would be. And nobody could rule out the possibility of some extreme climate shock, caused by processes perhaps not yet imagined in the convoluted systems that linked air, ice, seas, and living creatures.¹

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¹ E.g., “The threshold separating stable and unstable climate regimes represents a relatively small departure from the modern ice sheet configurations,” according to McManus et al. (1999), p. 1.